

SSEBop FANO Overview

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USGS EROS

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ET Estimation in Agriculture...

- It is a **RESPONSE** variable as opposed to precipitation (driver)
- It reflects the integrated effects of Energy/Aerodynamics, Soil Moisture, Vegetation and Environmental Stress

Potential

Energy

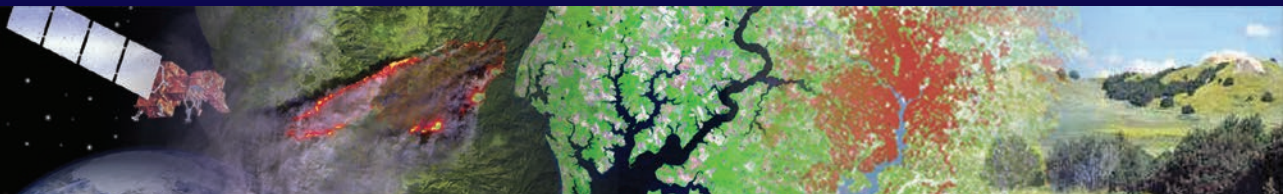
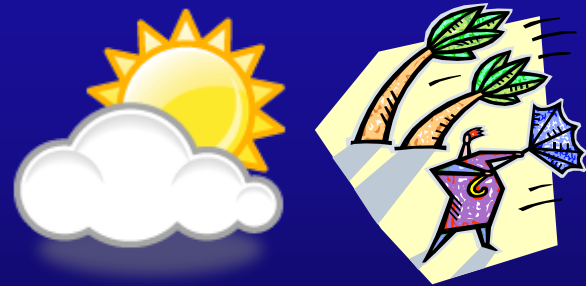
Wind/RH

Limitations

Moisture

Vegetation

Env. Stress



ET = Crop Water Use

Landsat 8



Field Flux Tower

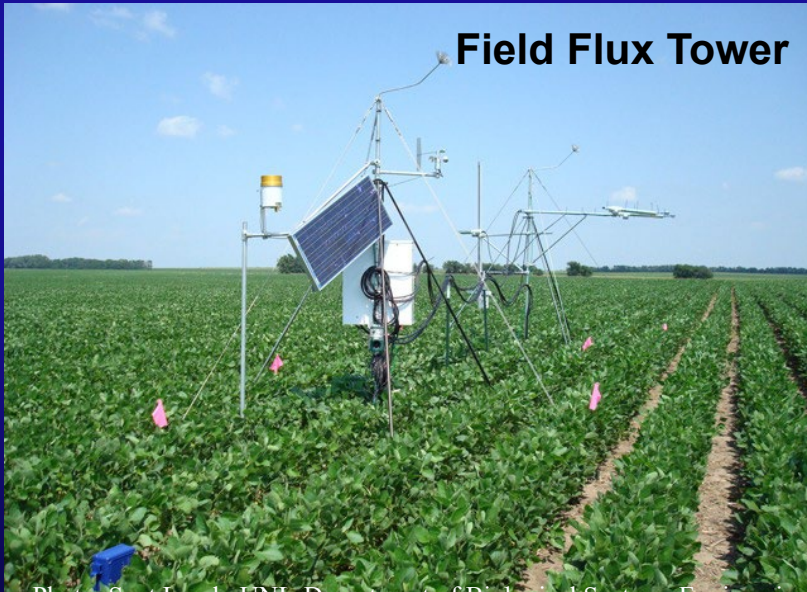
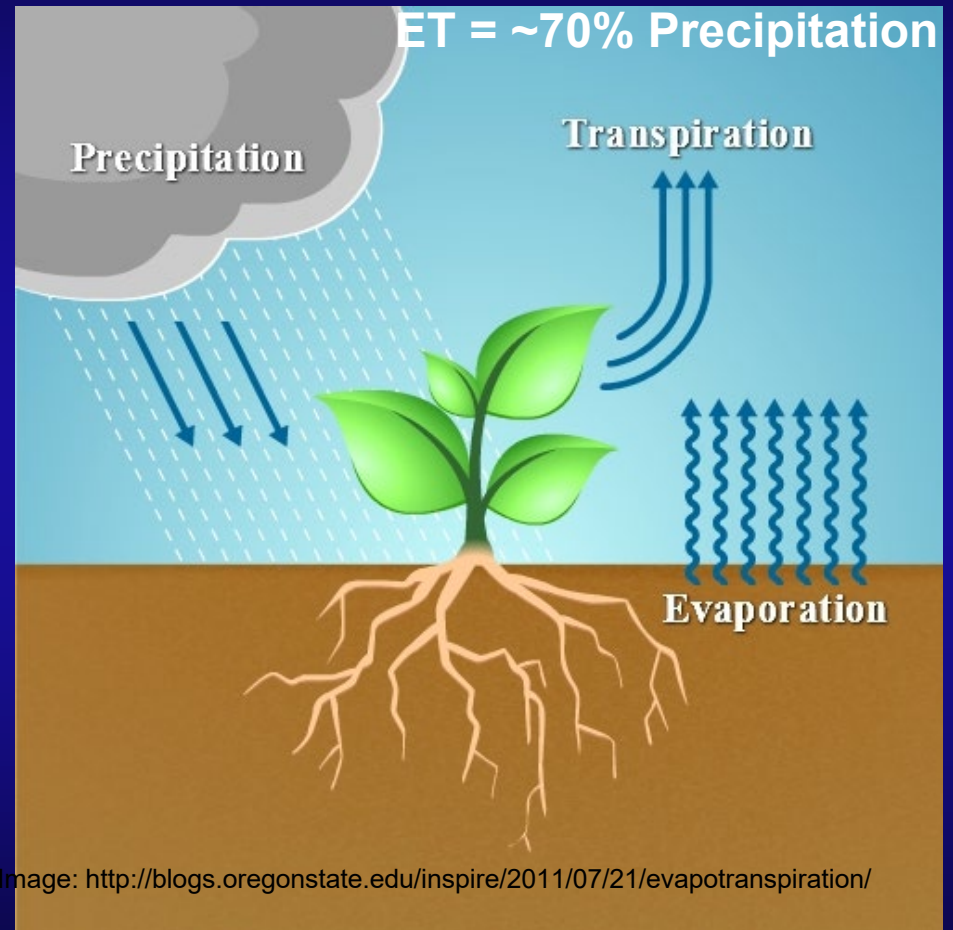
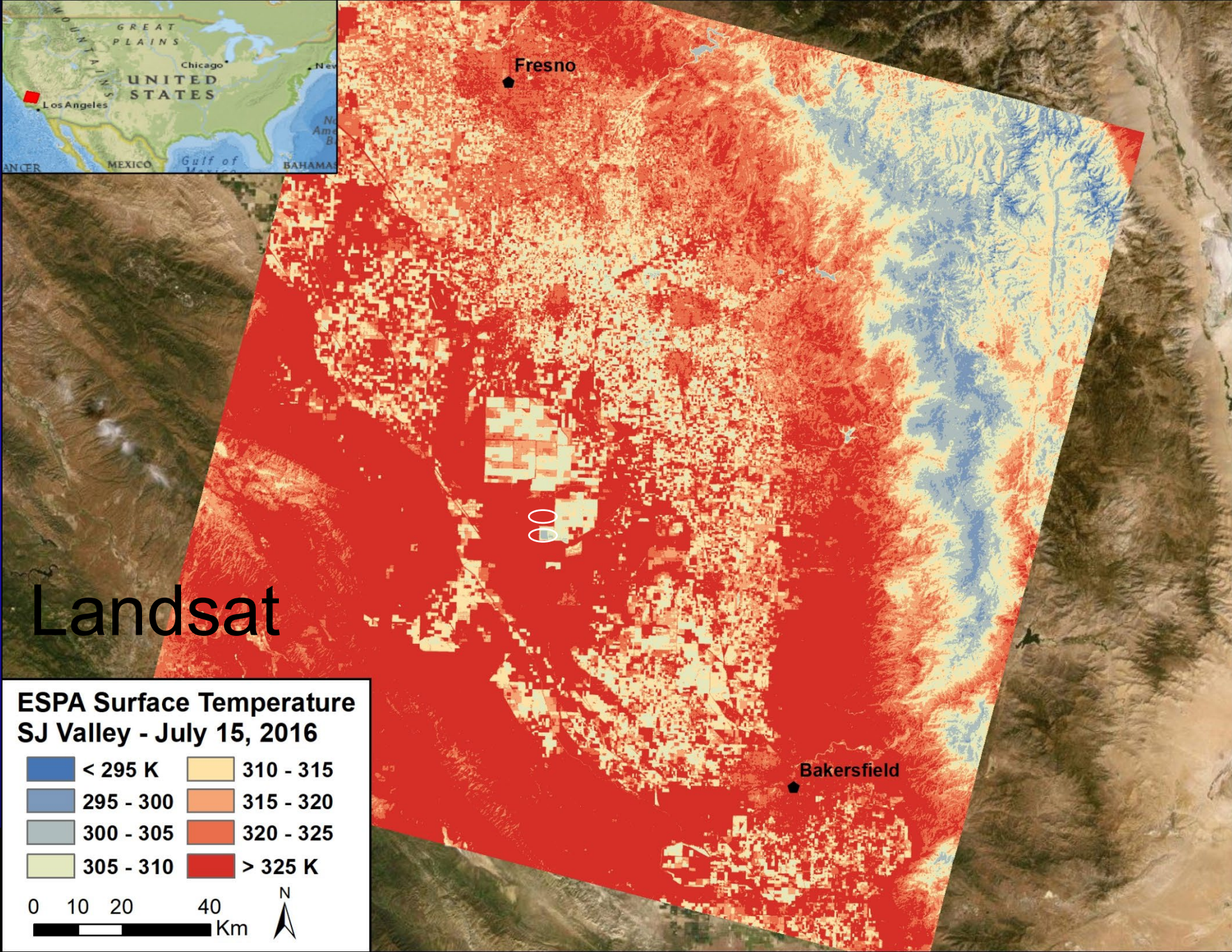


Photo: Suat Irmak, UNL-Department of Biological Systems Engineering

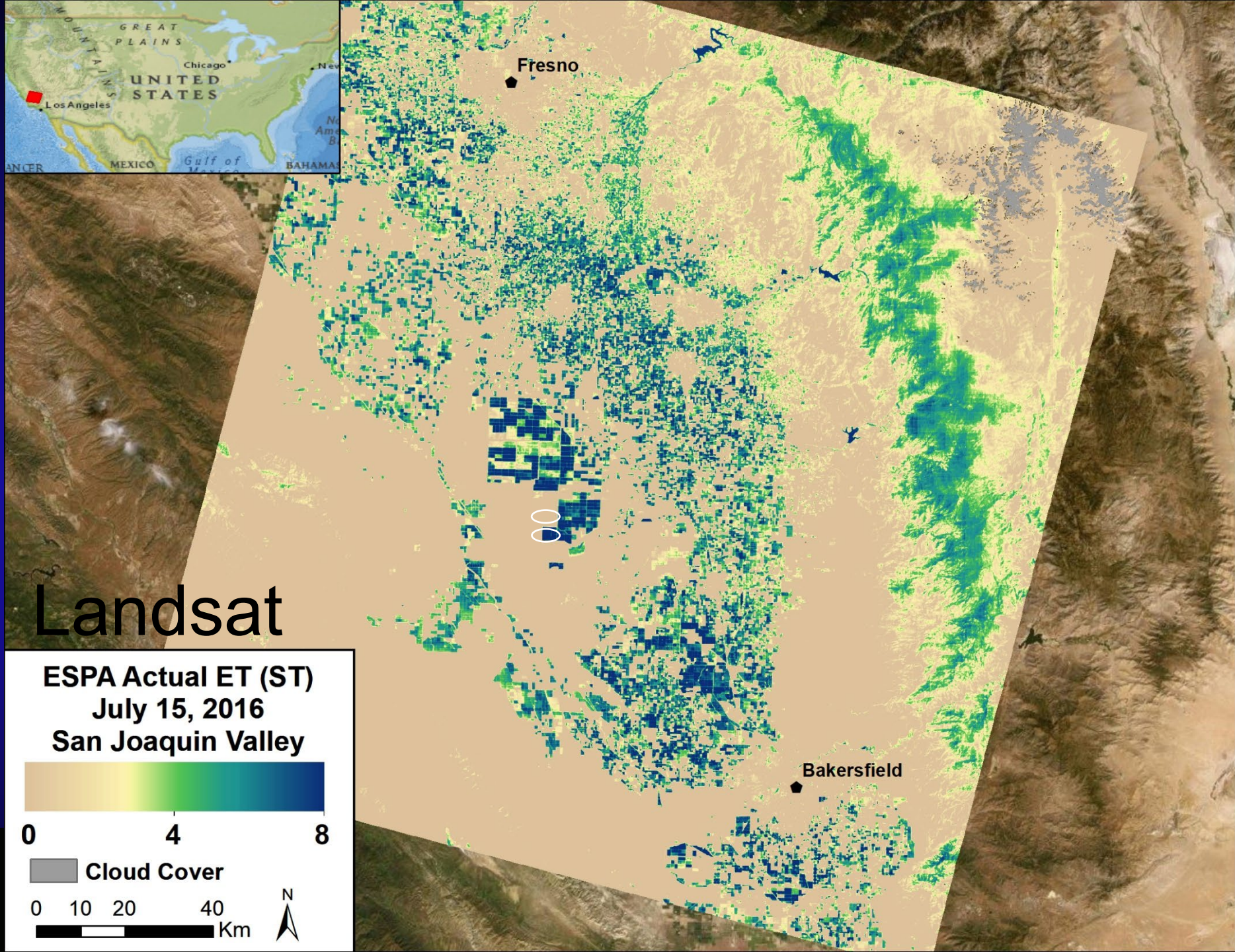
**Evapotranspiration =
transpiration + evaporation**





ESPA Surface Temperature SJ Valley - July 15, 2016





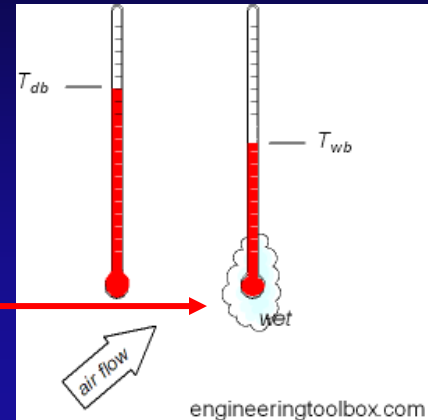
SSEBop Explained with “Psychrometry”

$$ea = es - \gamma (Td - Tw)$$

$$\gamma = \frac{C_p P}{\varepsilon \lambda} = 0.665 * 10^{-3} P$$

ea = actual vapor pressure (kPa)

es = saturated V.P. (kPa) at Tw



γ psychrometric constant [kPa °C⁻¹]

P atmospheric pressure [kPa],

λ latent heat of vaporization, 2.45 [MJ kg⁻¹]

c_p specific heat of air at constant pressure,
1.013 10⁻³ [MJ kg⁻¹ °C⁻¹]

ε ratio molecular weight of water vapor/dry air
= 0.622.

Ferrel, W.M. (1886); Allen et al. (1998)



Air vs “Surface” Psychrometry

(Thermometer vs Satellite Psychrometry)

Vapor pressure/relative humidity (standard psychrometry)

$$ea = es - \gamma (Td - Tw)$$

(Ferrel, W.M. (1886); Allen et al. (1998))

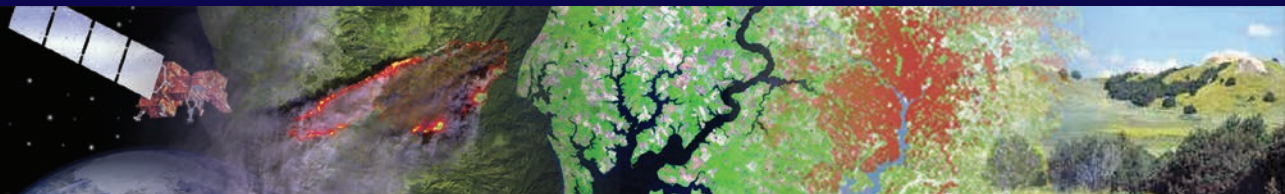
=>Large temp difference is a result of dry air (low RH) and hence low vapor pressure

Actual ET: Satellite “Psychrometry” Approach (SPA): SSEBop

$$ETa = ETo - \gamma^s (Ts - Tc)ETo$$

(Senay, 2018. App Eng. in Ag.)

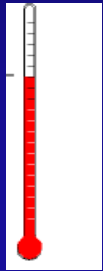
=>Large temperature difference is a result of dry soil/vegetation complex (low moisture, high stress) and hence low actual ET



Energy Balance Model: SSEBop



Land surface temperature differences are used to measure landscape water use rates through the effect of evaporative cooling

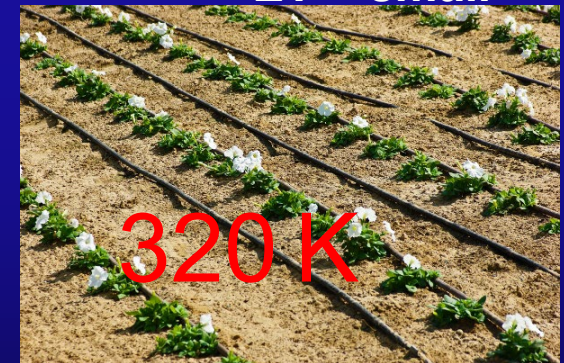


Surface Temp (T_s) 320 K

300 K
ET = Max

ET = moderate

Air Temp ($\sim T_c$) 300 K
ET = small

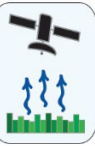


<http://www.croplife.com/equipment/irrigation-control-gets-ever-more-precise/>

<https://www.irrigationaustralia.com.au/about-us/types-of-irrigation/drip-irrigation>



Operational Simplified Surface Energy Balance (SSEBop)

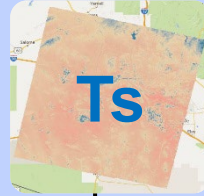


SSEBop
Evapotranspiration

parameters

NDVI, Air Temp, γ^s

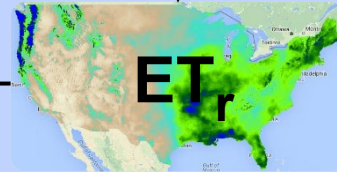
Surface Temp (T_s)



Weather Data

Radiation,
Temp, Wind,
RH, Pressure

ET_{fraction}



Reference ET

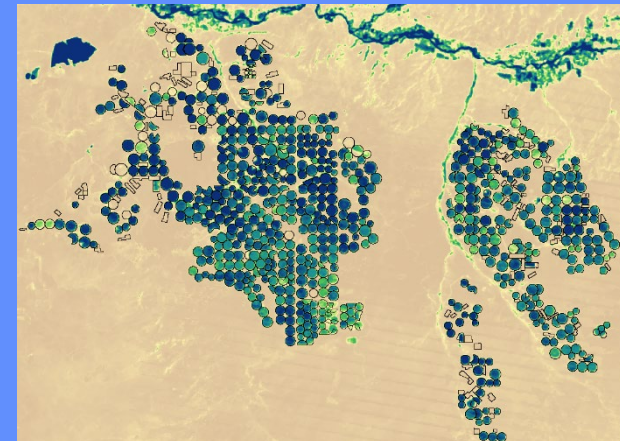
ET_a

$$ET_a = ET_f * ET_r$$

$$ET_f = 1 - \gamma^s(T_s - T_c)$$

$$\gamma^s = \frac{\rho \cdot C_p}{R_n \cdot r_{ah}}$$

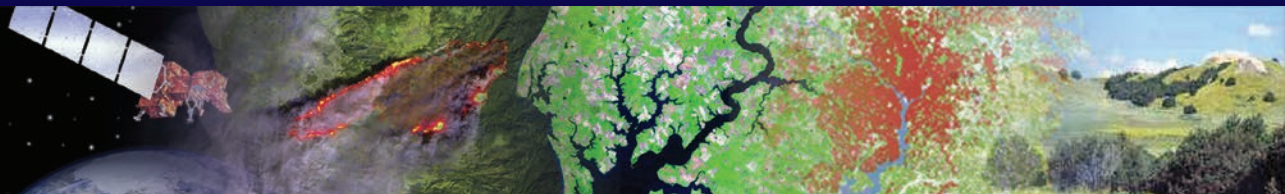
$$T_c = T_s - f \cdot dT(0.9 - NDVI): \text{FANO Equation (v6)}$$



Determining the Tc parameter

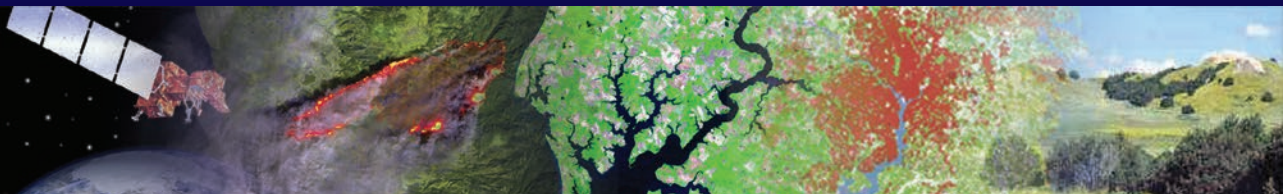
$$ET_f = 1 - \gamma^s(T_s - T_c)$$

- Satellite Psychrometry: the determination of the **wet-bulb**.
- Satellite acquires the ambient surface temperature (**dry-bulb**).
- We use the FANO Equation to determine Tc.



SSEBop: FANO

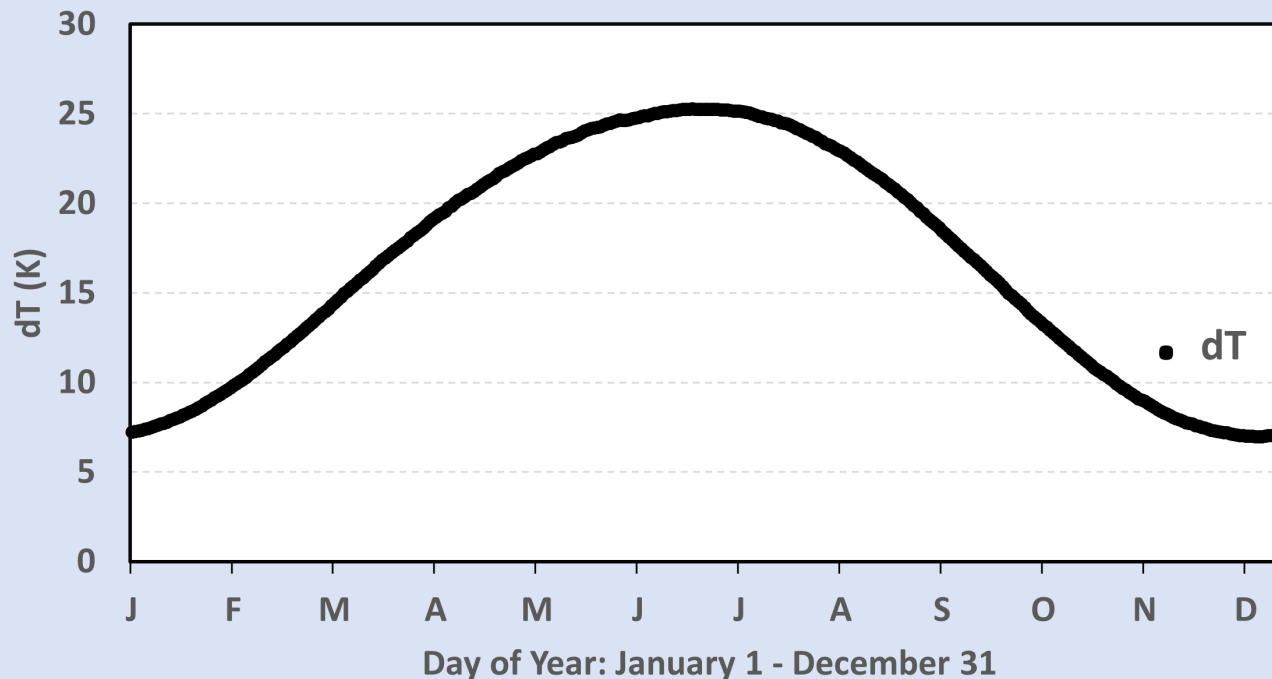
- A model parameterization approach using the **F**orcing **A**nd **N**ormalizing **O**peration.
- **Forcing**: it forces the landscape to have a wet-bulb value for the boundary condition. This represents the land surface temperature at maximum ET.
- **Normalizing**: parameters are normalized to formulate the linear FANO equation.



dT (1/ γ^s) parameter for SE Nevada

$$dT = \frac{R_n \cdot r_{ah}}{\rho \cdot C_p}$$

Input to the FANO Eqn.



dT : temperature difference [K, °C]
 R_n : Net radiation [$\text{MJ} \cdot \text{kg}^{-1} \cdot \text{d}^{-1}$]
 r_{ah} : aerodynamic resistance for heat [$\text{s} \cdot \text{m}^{-1}$]
 C_p : specific heat of air at constant pressure, $1.013 \cdot 10^{-3}$ [$\text{MJ} \cdot \text{kg}^{-1} \cdot \text{C}^{-1}$]

Kagone and Senay (2022)

Global dT is available at:

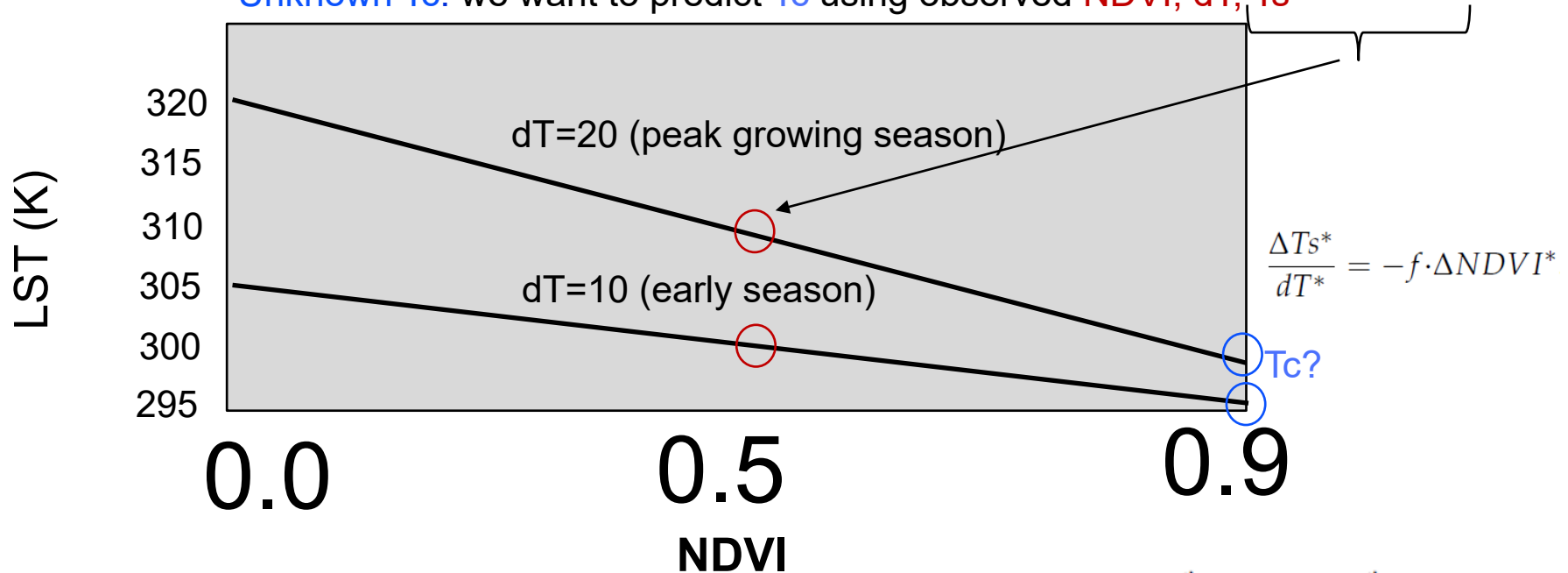
[Global gray-sky dT: the inverse of the surface psychrometric constant parameter in the SSEBop evapotranspiration model | U.S. Geological Survey \(usgs.gov\)](#)



Determining Tc using FANO Equation

Senay et al. (2023), RS.

- LST is inversely related to NDVI (“non-wet “ landscapes)
- But degree (slope) of relation varies by season
- Normalizing by dT accounts for seasonal variation
- Maximum NDVI with maximum ET ≈ 0.9 (assumed)
- **Unknown Tc**: we want to predict **Tc** using observed **NDVI**, **dT**, **Ts**



$$\Delta T_s^* = -f \cdot dT^* \cdot \Delta NDVI^*$$

$$\Delta NDVI^* = NDVI^* - NDVI_{max}$$

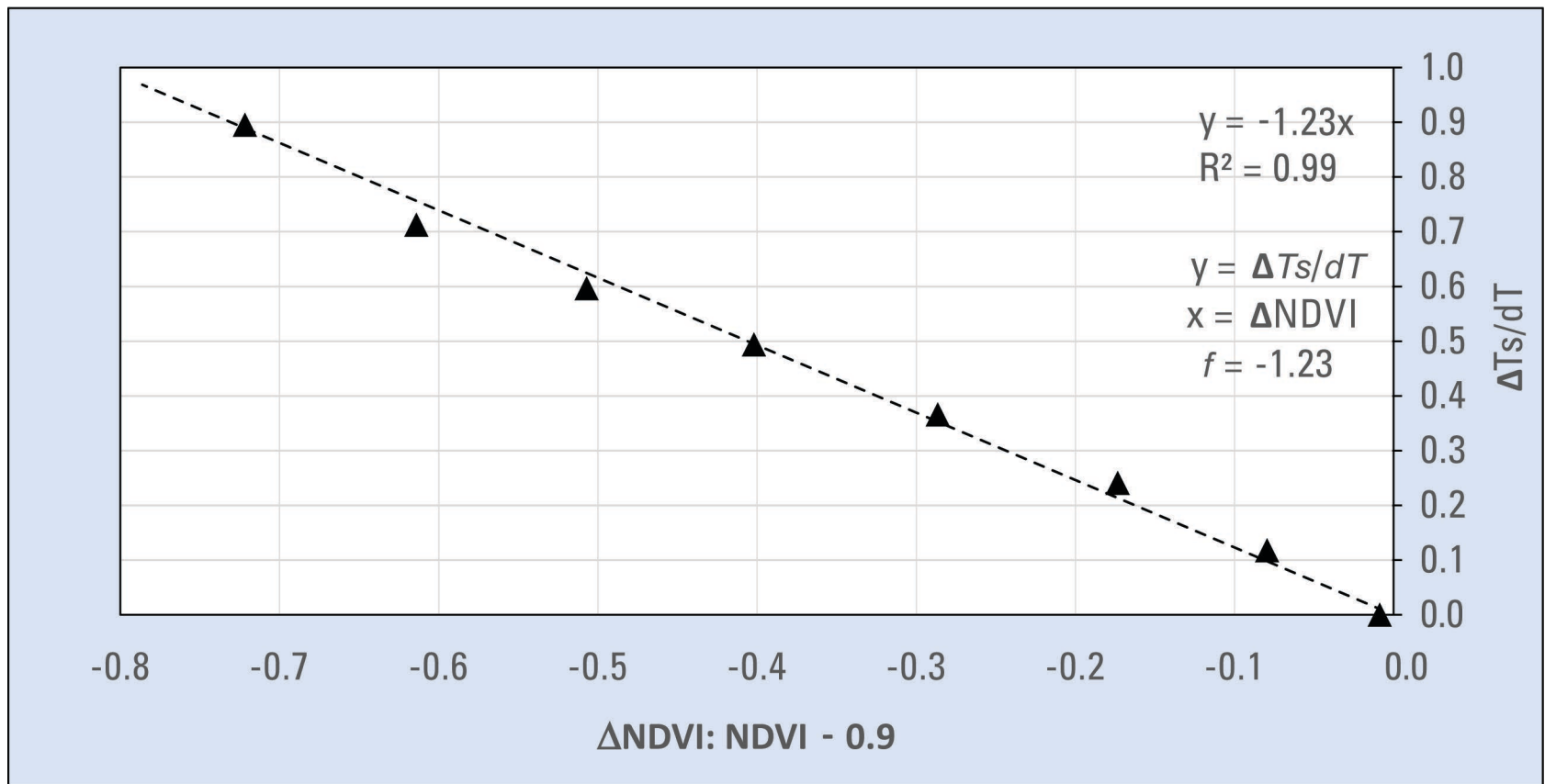
$$\Delta T_s^* = T_s^* - T_c^*$$

Same **NDVI** yields different **Tc** on different seasons.
Soln: Normalizing by dT



FANO Equation Formulation

$$\frac{\Delta T_s^*}{dT^*} = -f \cdot \Delta NDVI^* \quad \Rightarrow \quad T_c = T_s - f \cdot dT(0.9 - NDVI)$$



GET MAP LAYER

Variable ?

Type: Search Datasets

Climate & Hydrology

Dataset: ?

GridMET - 4km - Daily

Variable: ?

Precipitation (PPT)

Units: millimeters

Computation Resolution (Scale): ?

4000 m (1/24-deg)

Processing ?

Statistic (over day range): ?

Total

Calculation: ?

Values

Time Period ?

Period of Record: 1979-01-01 to 2023-08-04

Last 60 Days of Data

Start Date: 2023-06-06

End Date: 2023-08-04

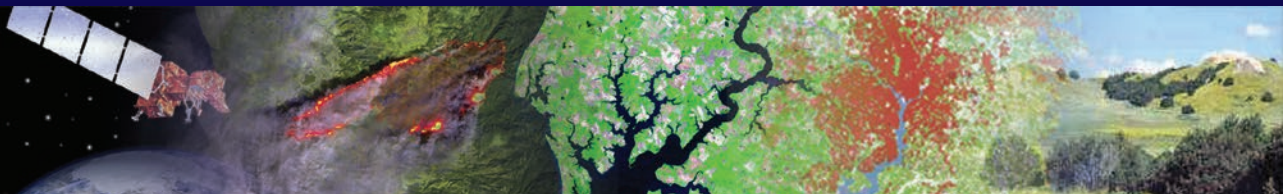
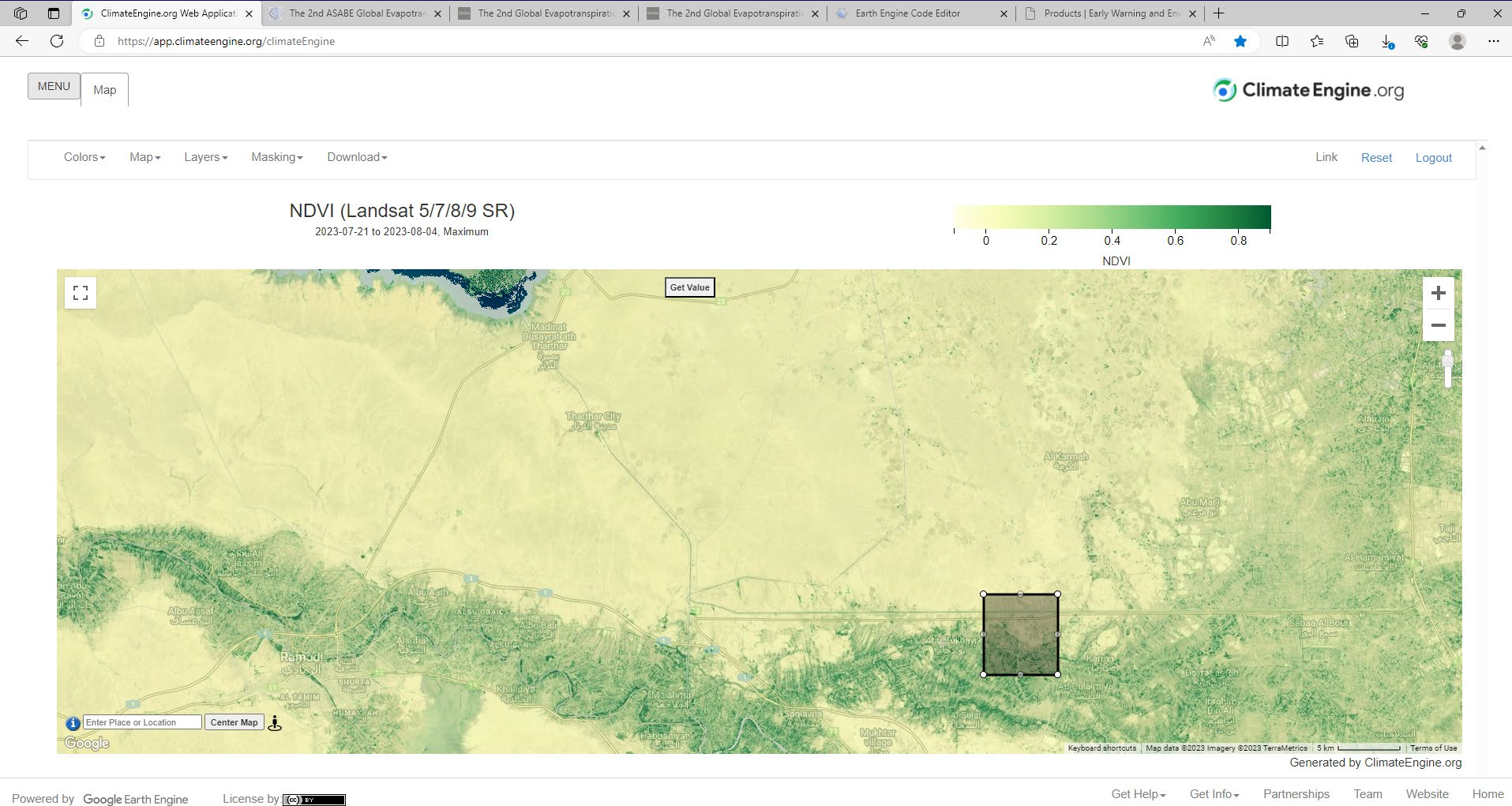
GET MAP LAYER

Colors | Map | Layers | Masking | Download

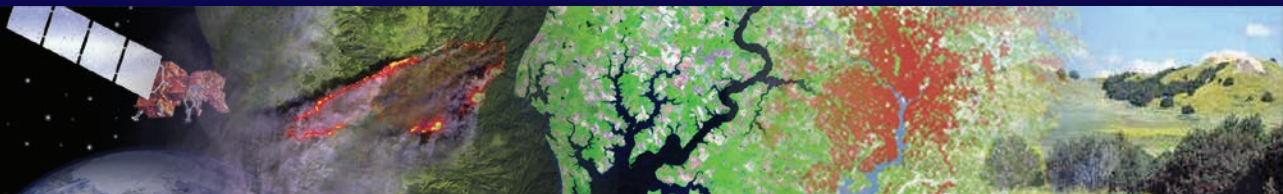
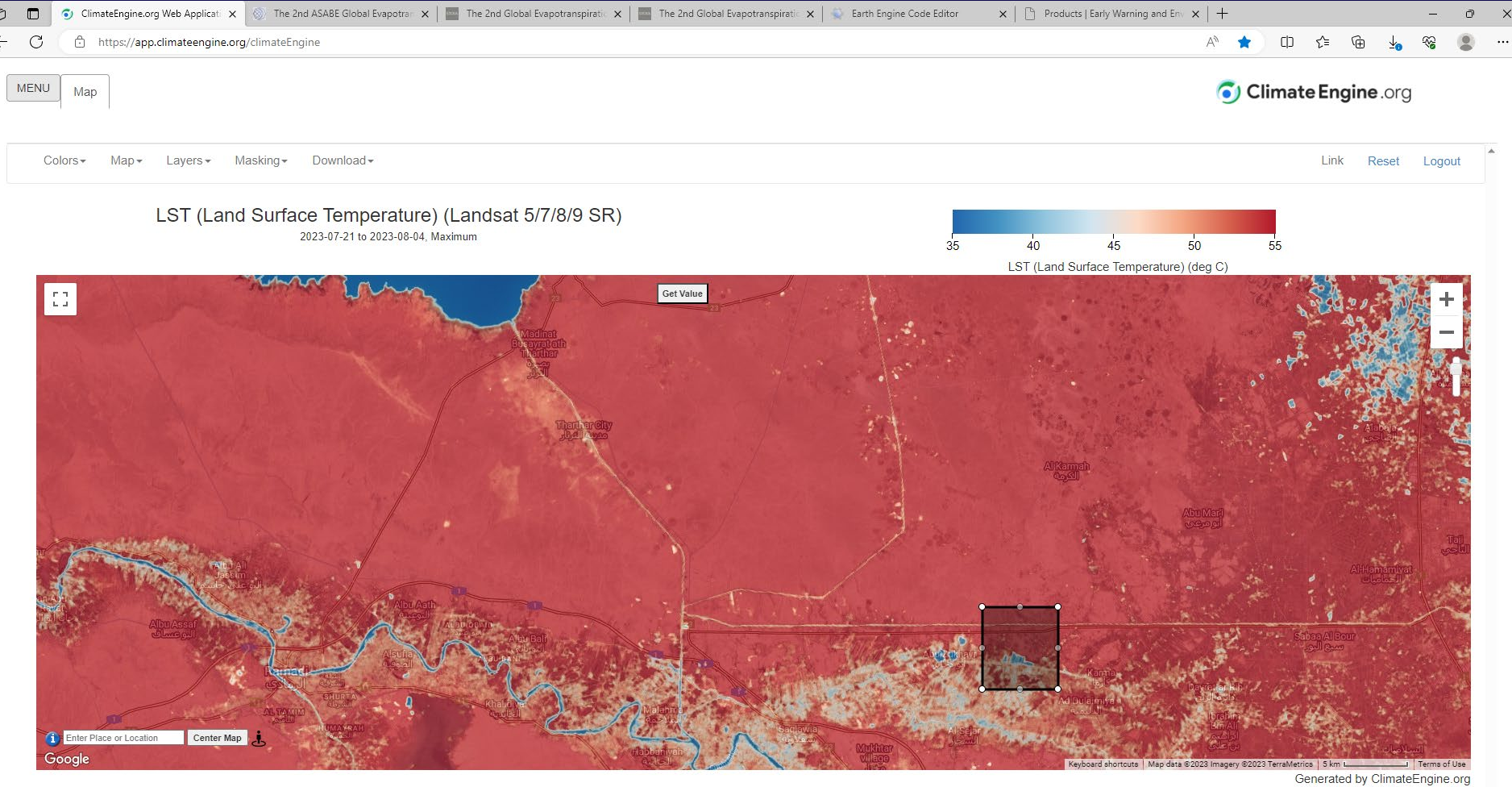
Link | Reset | Logout



NDVI 5-km = ~0.3



LST 5-km = $\sim 50 = 323$ K



Sample Calculation:

Assuming

$$dT^* = 20 \text{ K}$$

$$T_s^* = 323 \text{ K}$$

$$NDVI^* = 0.30$$

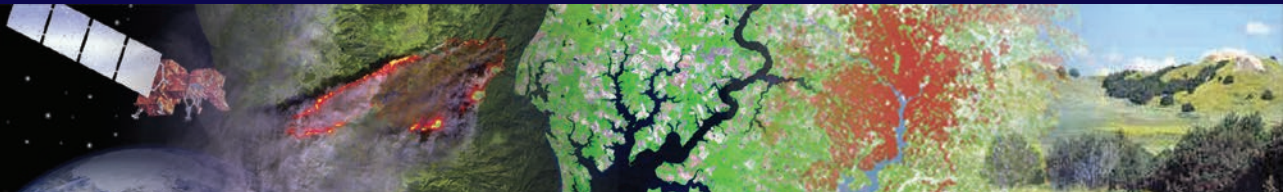
$$T_c^* = T_s^* - 1.25 * dT (0.9 - NDVI^*)$$

$$T_c^* = 323 - 1.25 * 20(0.9 - 0.3)$$

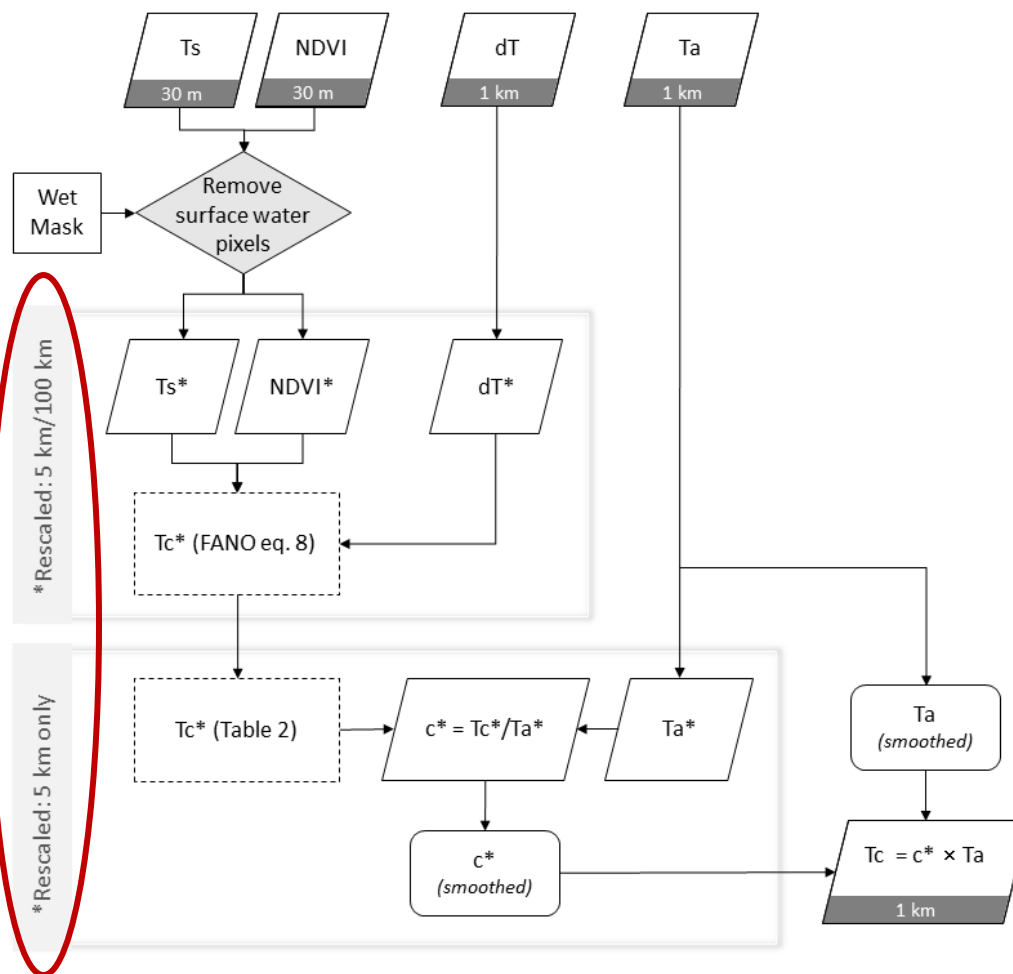
$$T_c^* = 308 \text{ K}$$

$$T_c^* = 35^\circ\text{C} \text{ (which is comparable to the water temp)}$$

TC* is then disaggregated to 1 km to Tc using air temperature (Ta). But Ta does not vary much over flat areas and its effect is limited.

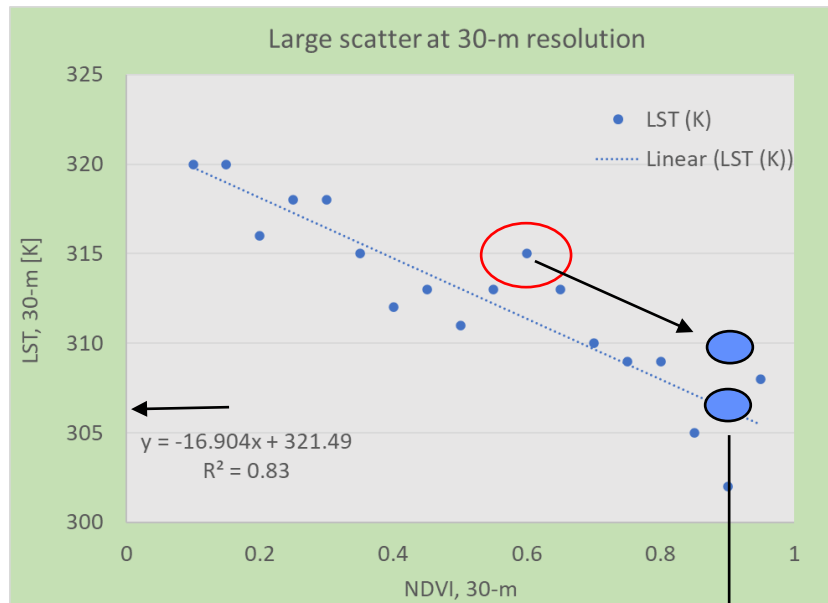


Tc Determination Steps

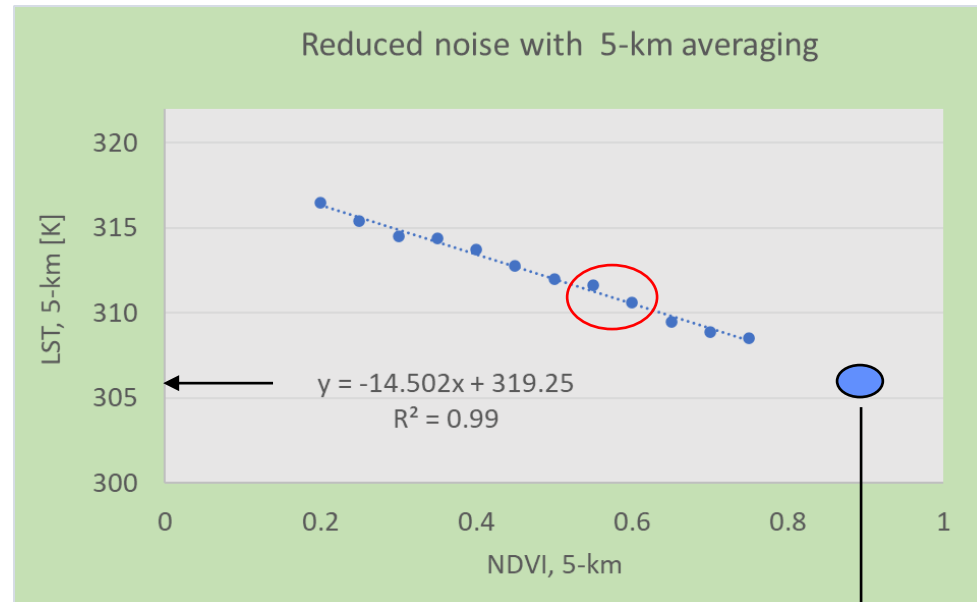


Why 5-km Spatial Averaging? To Reduce Noise

Avoid over- or under-predication of T_c for **neighboring pixels**.



NDVI = 0.90
LST = 306.3 K

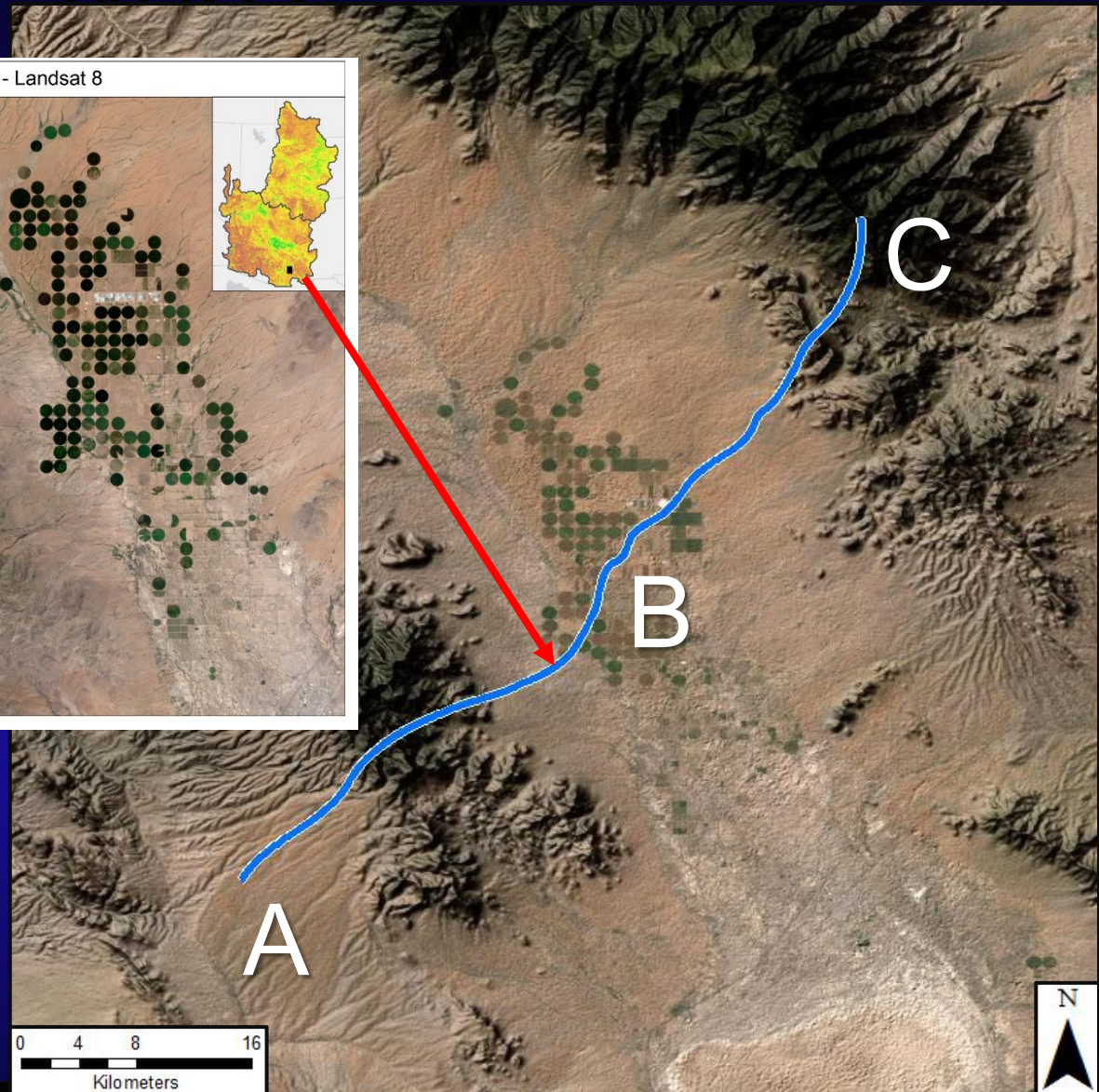
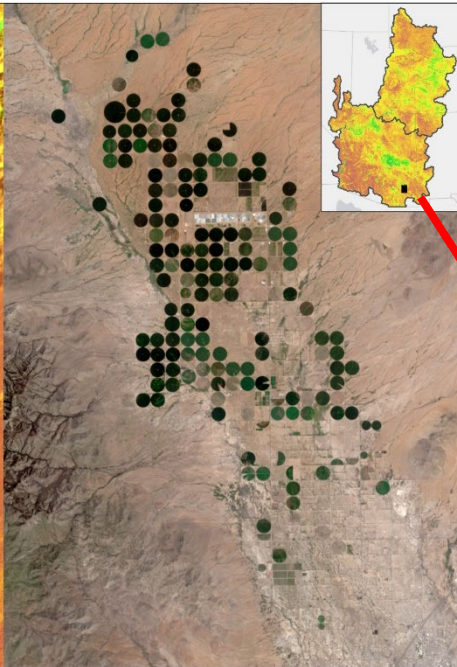
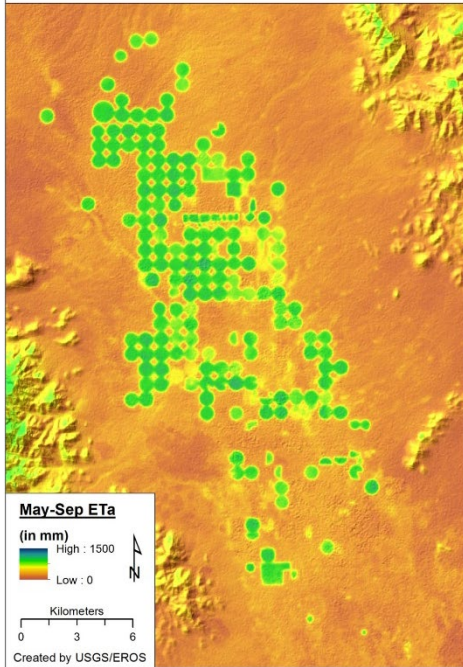


NDVI = 0.90
LST = 306.2



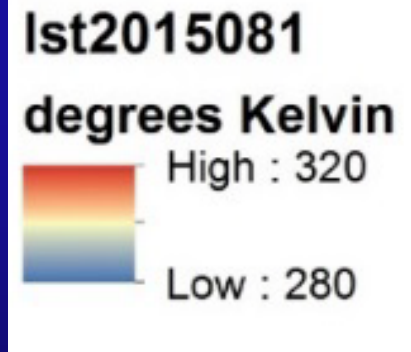
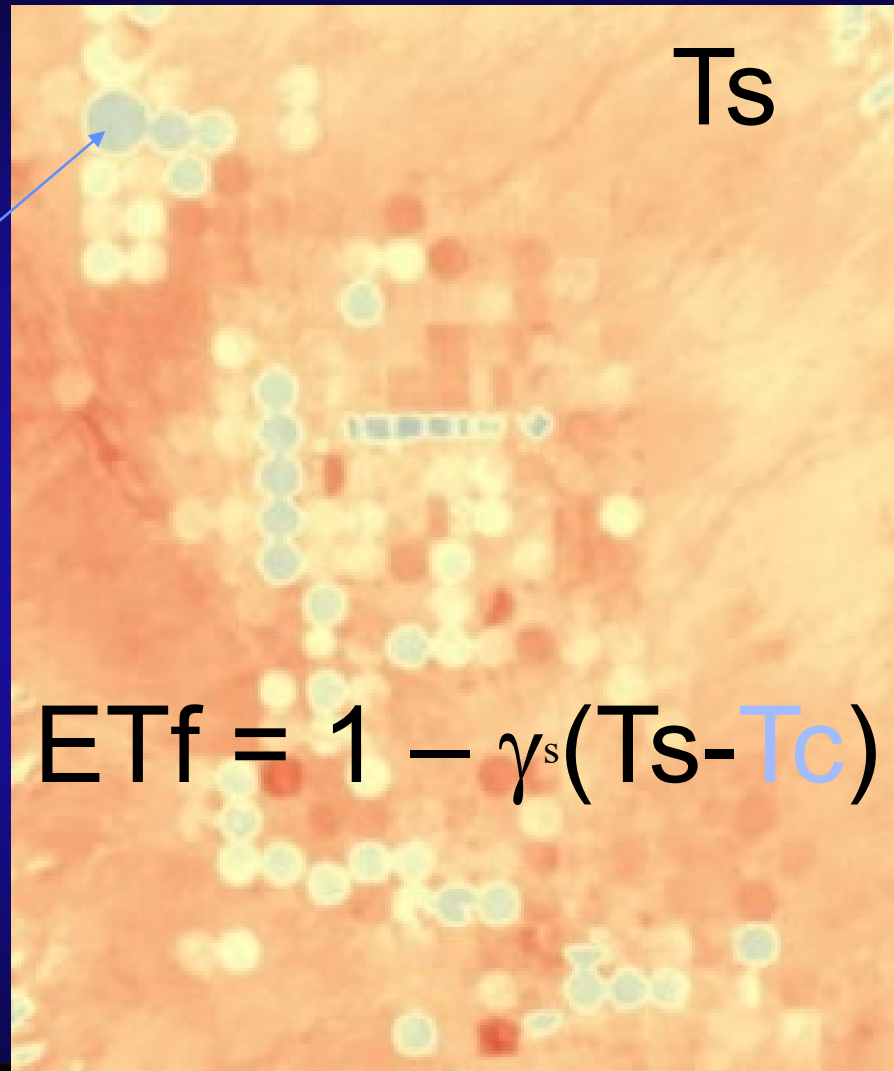
Transect in Willcox Irrigation Basin

Willcox Basin Irrigation - 2013 - Landsat 8



LST (Ts) distribution in an irrigated valley Willcox, AZ,: March 21, 2015)

$T_s = T_c$

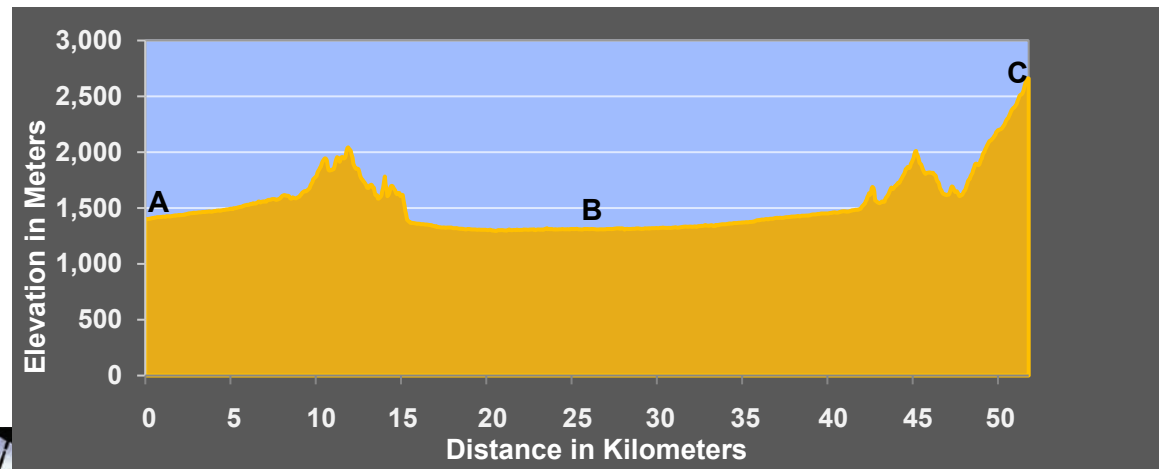
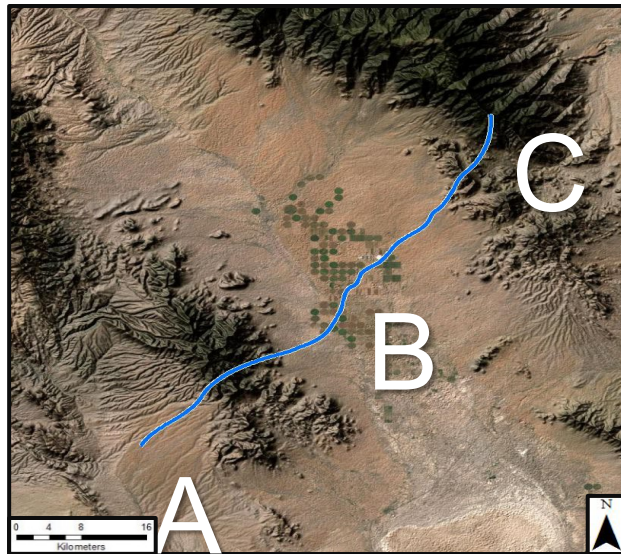
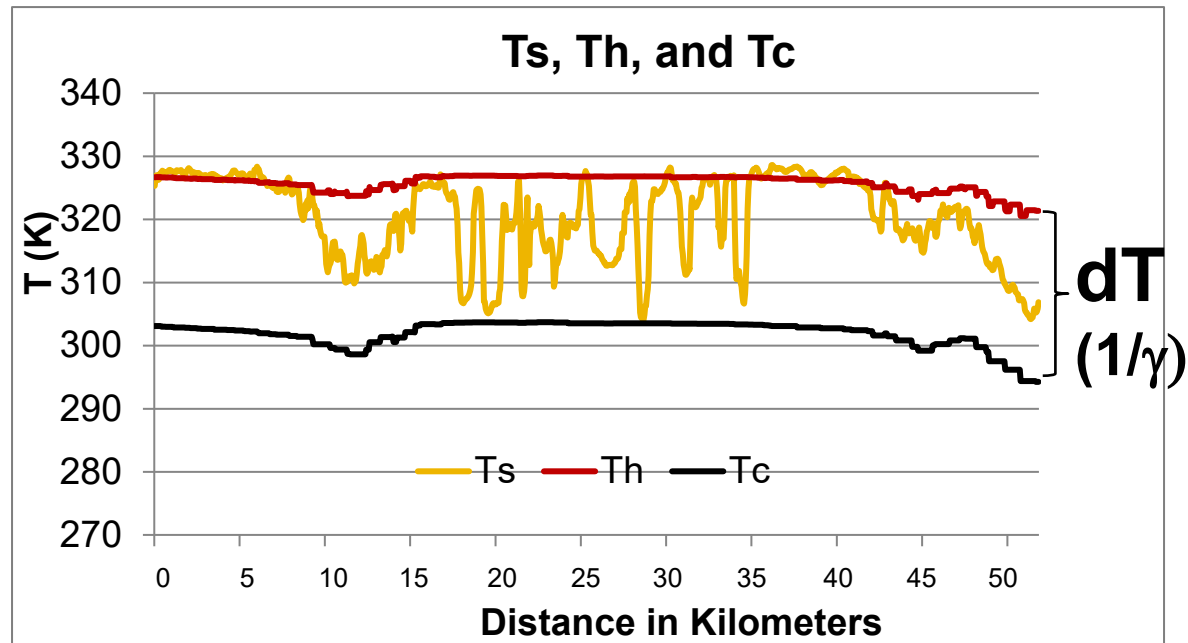


LST(Ts), Hot (Th)/Cold (Tc) Limits for Jun 23, 2014

$$ETf = 1 - \gamma^s (Ts - Tc) \quad (0 - 1)$$

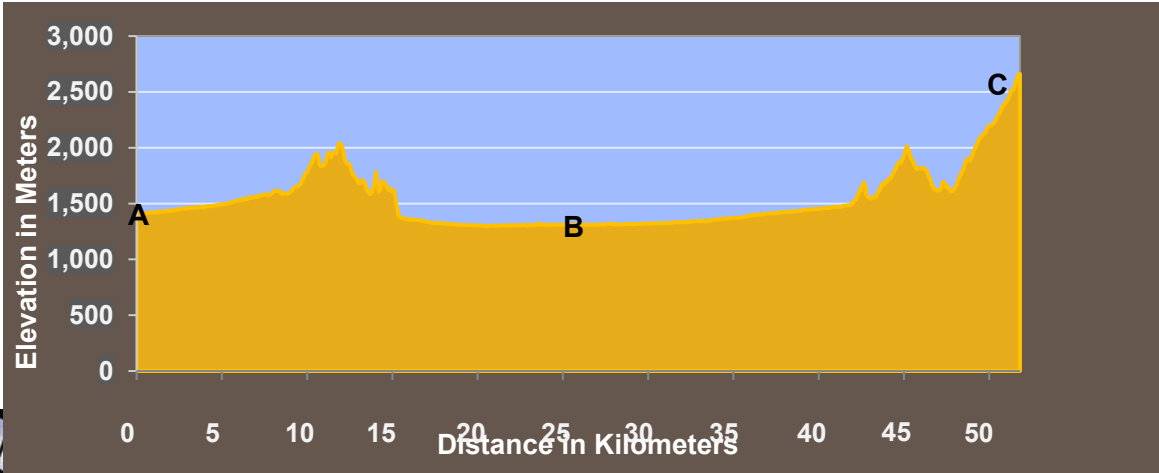
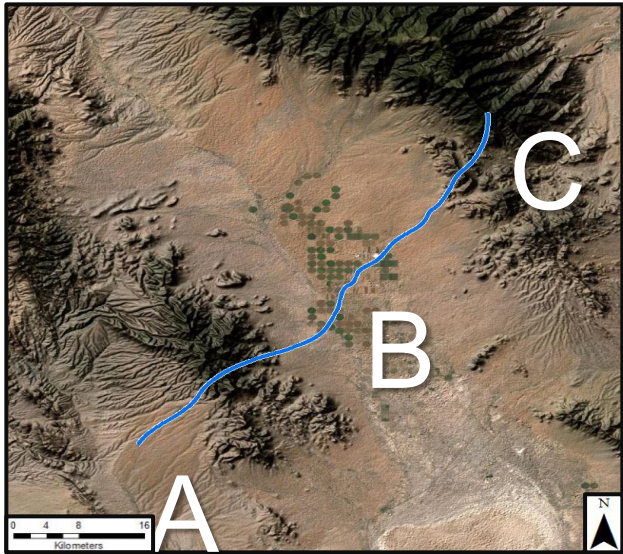
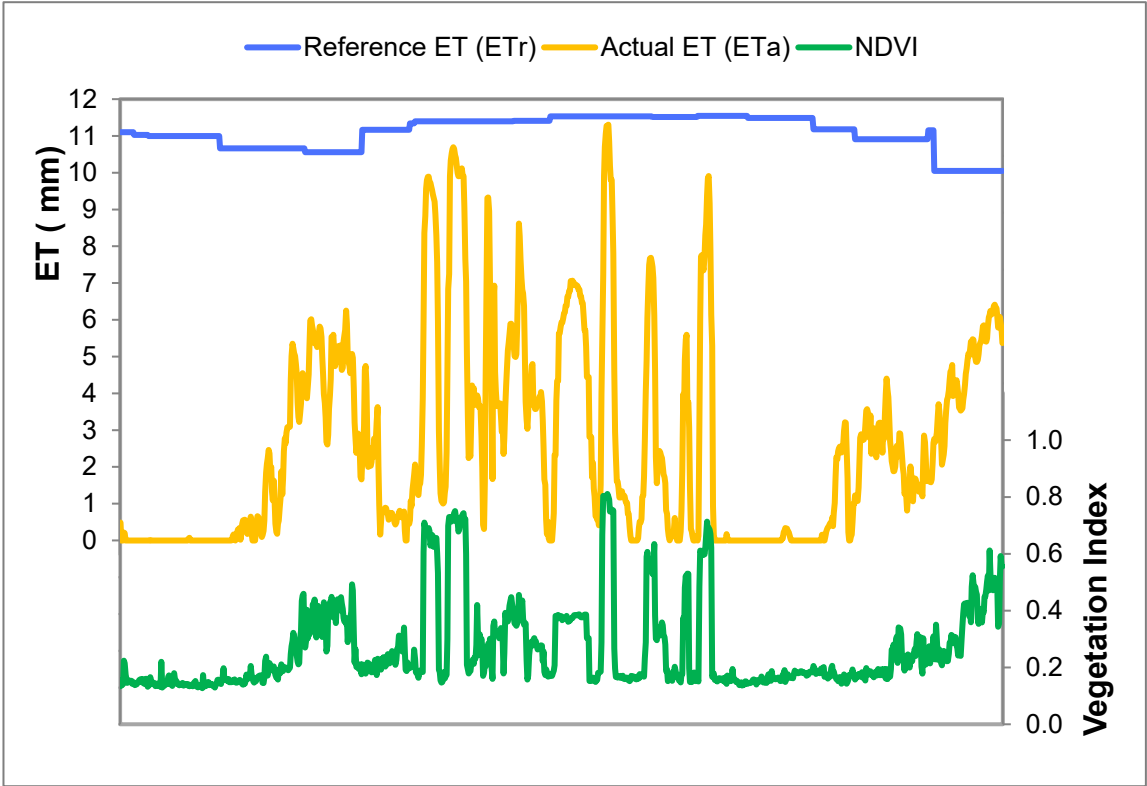
$$\gamma^s = \frac{\rho C p}{R n \gamma_a}$$

$$ETa = ETo - \gamma^s (Ts - Tc) ETo$$

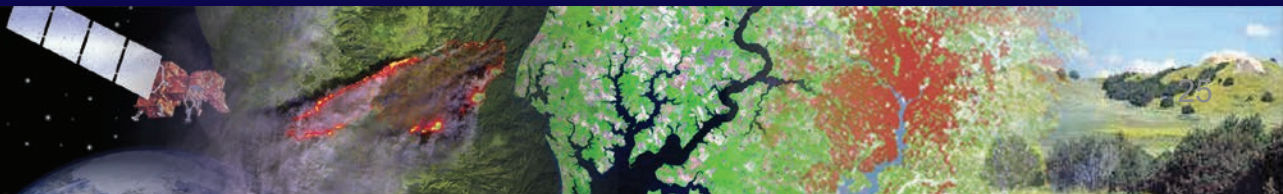
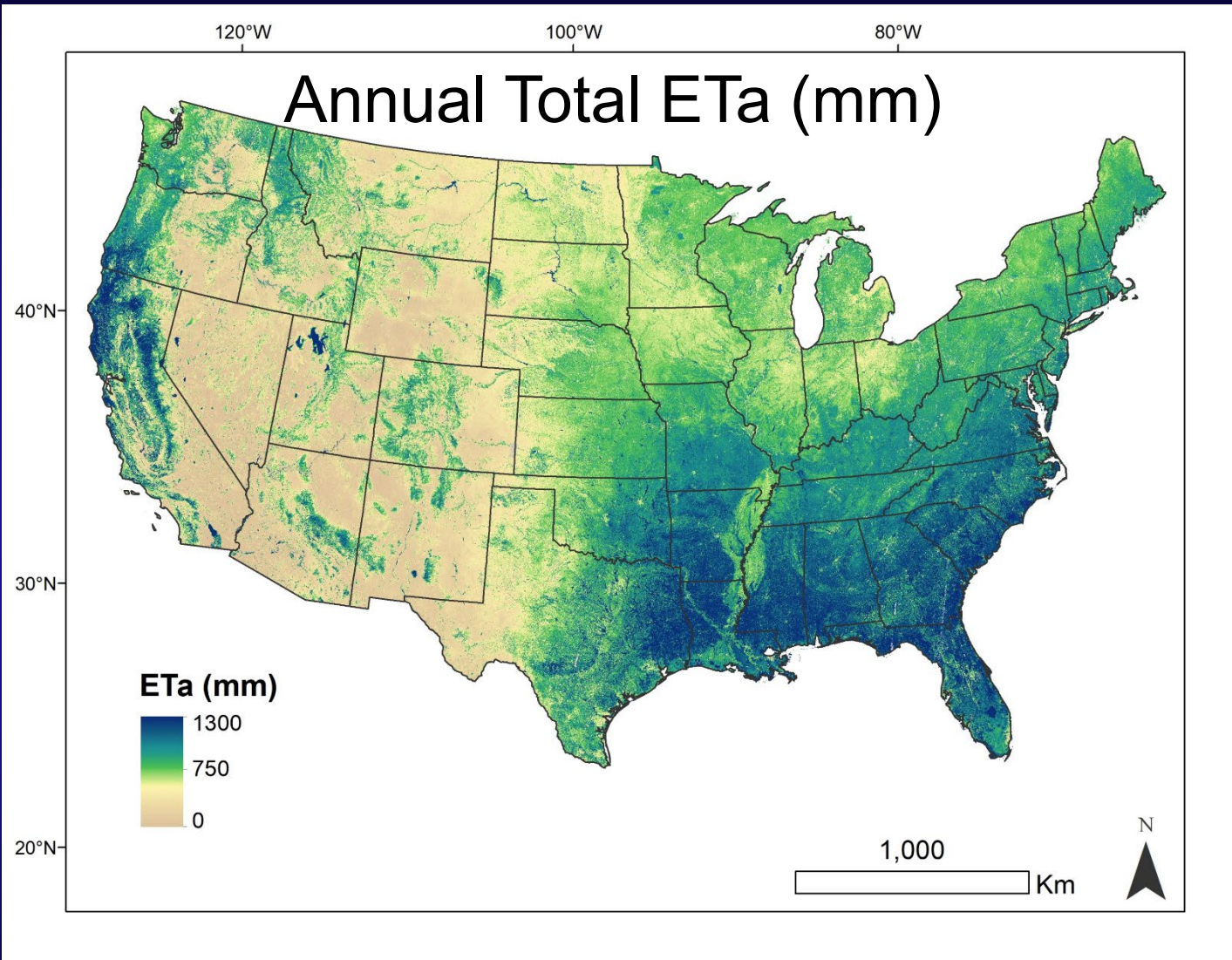


ETa, PET (ETr), NDVI

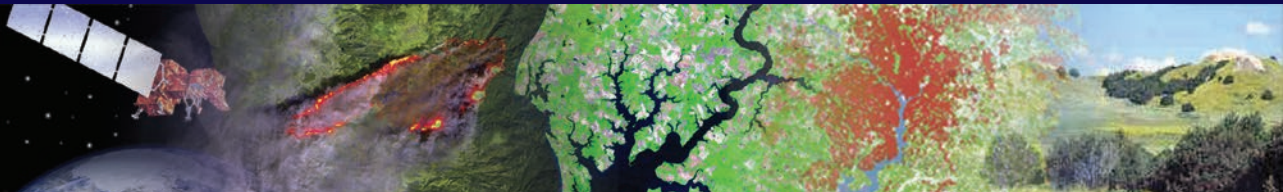
$$ETa = ETo - \gamma^s (Ts - Tc) ETo$$



1st Ever Landsat Scale CONUS ETa



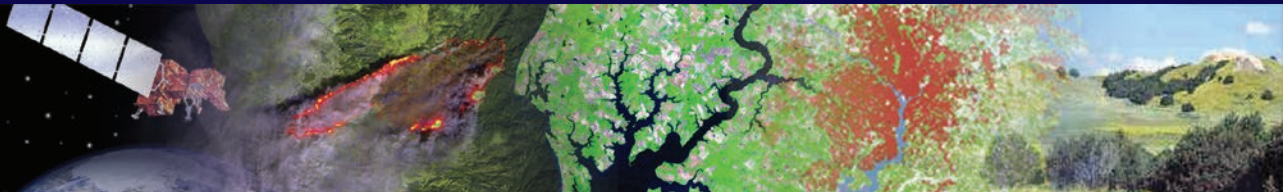
Questions?



Challenge: ET under potential vs water limiting conditions

- Landscape is at different levels of stress; thus, **actual ET** is \leq **potential**.
- Allen et al (1998)
 - **ET = K_s * K_c * E_{To}**
 - K_c = **type and stage** of crop (~0.15 – 1.0)
 - K_s = **soil moisture** stress factor (0 to 1)

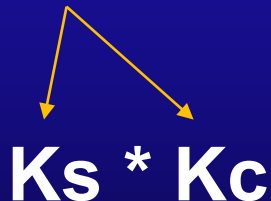
(This requires knowledge of crop types, stage and moisture distribution)



More direct estimation of stress using remote sensing approaches...

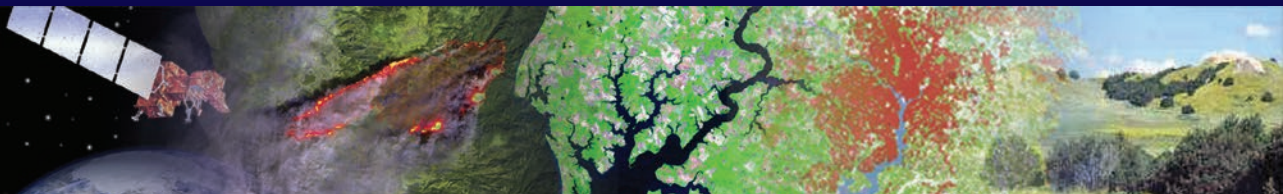
Land surface temperature (LST) derived from remotely sensed imagery can be used to estimate the combined effects of soil moisture and environmental stress factors on vegetation.

$$ET = ET_f * ET_o$$



A diagram with two orange arrows pointing from the term ET_f in the equation above to the term $K_s * K_c$ below it.

$$K_s * K_c$$



Energy Balance Approach for ET:

Accounts for water, agronomic and environmental stresses

USGS **WaterSMART** and **FEWS NET** use the **SSEBop** (Operational Simplified Surface Energy Balance) approach for:

- 1) Water Use and Availability Assessment
- 2) Drought Monitoring & Early Warning

